

Glacial deposits provide a long-term record of climate and sea level changes on Earth. Detailed study of sedimentary rocks deposited during and immediately after glacial episodes is paramount to accurate paleoclimatic reconstructions and for our understanding of global climatic and eustatic changes. This book presents new information and interpretations of the ancient glacial record, looking in particular at the Late Proterozoic and Late Paleozoic eras. The influence of global tectonics on the origins and distribution of ice masses and the character of glacial deposits through geologic time is emphasised. Sequence stratigraphic techniques are applied to glaciogenic successions, and explanations are put forward for possible low-latitude glaciation during the Late Proterozoic era and the association of carbonate deposits with glaciogenic rocks. Early interglacial conditions, represented by dark-grey mudrocks and ice keel scour features, are discussed. These studies, from key workers in International Geological Correlation Program Project 260, will aid the understanding of Earth's climatic history.



Earth's Glacial Record



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Earth's Glacial Record

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Contents

/ 1 (1 1 1 1	Preface Acknowledgements Geodynamic controls on glaciation in Earth history N. EYLES and G.M. YOUNG Abstract Introduction Late Archean—Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch Neoproterozoic glaciations	xiii xvii 1 1 1 2		northern Utah (formation of Perry Canyon) Sheeprock Mountains and Deep Creek Range, western Utah (Sheeprock Group and Horse Canyon Formation) Southeastern Idaho (Pocatello Formation) Interpretation of facies associations	37 37 38
1 (1 1 1 1	Geodynamic controls on glaciation in Earth history N. EYLES and G.M. YOUNG Abstract Introduction Late Archean–Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	1 1 1 2		western Utah (Sheeprock Group and Horse Canyon Formation) Southeastern Idaho (Pocatello Formation)	
1 (1 1 1 1	Geodynamic controls on glaciation in Earth history N. EYLES and G.M. YOUNG Abstract Introduction Late Archean–Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	1 1 1 2		Canyon Formation) Southeastern Idaho (Pocatello Formation)	
1 1 1 1	N. EYLES and G.M. YOUNG Abstract Introduction Late Archean–Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	1 1 2		Canyon Formation) Southeastern Idaho (Pocatello Formation)	
1 1 1 1	N. EYLES and G.M. YOUNG Abstract Introduction Late Archean–Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	1 1 2		Southeastern Idaho (Pocatello Formation)	38
I I	Abstract Introduction Late Archean–Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	1 2			
I	Introduction Late Archean–Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	1 2		THIEIDIEIALION OF IACIES ASSOCIATIONS	38
I	Late Archean-Paleoproterozoic glaciations Mesoproterozoic non-glacial epoch	2		Massive diamictite facies association: ice-proximal	
ľ	Mesoproterozoic non-glacial epoch			non-reworked deposits	38
		3		Stratified diamictite and graded sandstone	
1	Neophoterozoic giaciations	5		association: sediment gravity flow deposits	39
	Overview	5		Diamictite and laminated sandstone association: ice-	
		6		proximal deposits with abundant meltwater	39
	Timing and tectonic setting	7		Carbonate, shale and sandstone association: non-	57
	The importance of extensional rift basins	7		glacial marine deposits	40
	South Australia/northern Canadian Cordillera	8		Cross-bedded sandstone association: fluvial deposits	40
	North Atlantic sector			Discussion: facies distribution, subsidence mechanisms,	40
	Discussion	9		uncomformities and preservation potential in rift	
J	Early and Late Paleozoic glaciations	10			40
	South America	11		settings	40 40
	South Africa	14		Subsidence mechanisms	
	Australia	15		Relation of facies to tectonic position	41
	Antarctica	17		Generation of unconformities	41
	Southeast Asia	17		Geologic setting of the Rapitan glaciation	42
	Arabian Peninsula	17		Acknowledgements	42
	India and Pakistan	17		References	42
	Discussion	18	3	The Neoproterozoic Konnarock Formation,	47
	Mesozoic glaciation	18		southwestern Virginia, USA: glaciolacustrine facies in a	
]	Late Cenozoic glaciations	19		continental rift	
	The role of tectonic uplift	19		J.M.G. MILLER	
	The role of ocean currents	21		Abstract	47
(Conclusions	22		Introduction	47
1	Acknowledgements	22		Stratigraphic and structural setting	48
]	References	22		Paleotectonic setting	50
				Facies and facies associations	50
	Glacial-marine facies in a continental rift environment:	29		Argillite	50
]	Neoproterozoic rocks of the western United States			Sandstone	53
(Cordillera			Diamictite	54
]	P.K. LINK, J.M.G. MILLER and N. CHRISTIE-BLICK			Environmental synthesis	55
	Abstract	29		Conclusion	57
]	Introduction	29			57
	Evidence for overall glacial origin	31		Acknowledgements	57 57
	Facies associations	31		References	31
	Review of stratigraphy	34	4	Glaciogenic deposits of the Permo-Carboniferous	60
	Death Valley region of southeastern California			Dwyka Group in the eastern region of the Karoo Basin,	
	(Kingston Peak Formation)	36		South Africa	
	Central Wasatch Range and Antelope Island,			V. VON BRUNN	
	northern Utah (Mineral Fork Formation)	36		Abstract	60

vii



Frontmatter/Prelims

More information

viii Contents

	Introduction	60		Poorly stratified diamictite facies	103
	Pre-Dwyka highland surface (Northern Region)	60		Bedded diamictite facies	103
	Zone of irregular palaeorelief (Intermediate Region) Diamictite–mudrock facies association	62 63		Dropstone-bearing laminated lithofacies Thin-bedded to banded-laminated siltstone-mudstone	103
	Sandstone-arenaceous diamictite-conglomerate			facies	104
	facies association	63		The Wudaotang Section in Helan County	104
	Pre-Dwyka lowland plain (Southern Region)	65		The Tuerkeng Section	105
	Sequence of events	66		The Zhengmuguan Section	105
	Discussion	67		The Jingdiquan Section	105
	Acknowledgements	68		REE geochemical characteristics of the Zhengmuguan	100
	References	68		Formation	106
5	Itararé Group: Gondwanan Carboniferous-Permian of	70		Sedimentary facies and environmental analysis Conclusions	107 107
	the Paraná Basin, Brazil			References	107
	A.B. FRANÇA				
	Abstract	70	9	Architectural styles of glacially influenced marine	109
	Introduction	70		deposits on tectonically active and passive margins	
	Tectonic setting	70		M.R. GIPP	
	Stratigraphy	71		Abstract	109
	Lagoa Azul Formation	72 72		Introduction The Gulf of Alaska	109
	Campo Mourão Formation	73			109
	Taciba Formation	75 76		The Scotian margin Models	110
	Summary and conclusions	76 82		Passive margin	111 111
	Acknowledgements References	82 82		Active margin	111
	References	02		Similarities	113
6	The interpretation of massive rain-out and debris-flow	83		Distinguishing features in the models	116
	diamictites from the glacial marine environment			Other examples	117
	J.N.J. VISSER			Summary	119
	Abstract	83		Acknowledgements	119
	Introduction	83		References	119
	Massive diamictites from the Dwyka Formation	84	10	Maning Assurance and a second	
	Floriskraal	84	10	Marine to non-marine sequence architecture of an	121
	Elandsvlei	86		intracratonic glacially related basin. Late Proterozoic of the West African platform in western Mali	
	Kransgat River	87		J.N. PROUST and M. DEYNOUX	
	Douglas	90		Abstract	121
	Discussion	92		Introduction	121
	Acknowledgements	94		Baselevel concept	121
	References	94		Geological setting	122
7	Neoproterozoic tillite and tilloid in the Aksu area,	95		Definition of the depositional genetic unit from the	
	Tarim Basin, Uygur Xinjiang Autonomous Region,			sediment architecture of a key exposure	124
	Northwest China			Facies distribution in architectural elements	127
	LU SONGNIAN and GAO ZHENJIA			Landward-stepping unit	127
	Abstract	95		Vertical stacking unit	130
	Introduction	95		Seaward-stepping unit	132
	The Precambrian stratigraphic sequences	95		Discussion	135
	Sedimentary features of diamictites in the Qiaoenbulak			Bimodal facies architecture and time/space	
	Group	97		distribution of sediments and bounding surfaces	
	Features of the diamictite of the Umainak Formation			within genetic units	135
	(Wushinanshan Group)	98		Stacking cycle	137
	Comparison between diamictites of the Qiaoenbulak	00		Comparison with genetic units described in the	
	Group and the Wushinanshan Group Acknowledgements	98		literature	138
	References	99 99		Estimated baselevel cycles duration	139
	References	99		Inferred factors controlling sedimentation	139
8	Lithology, sedimentology and genesis of the	101		Conclusion	141
	Zhengmuguan Formation of Ningxia, China			Acknowledgements	141
	ZHENG ZHAOCHANG, LI YUZHEN, LU SONGNIAN			References	142
	and LI HUAIKUN		11	The enigmatic Late Proterozoic glacial climate: an	146
	Abstract	101		Australian perspective	
	Introduction	101		G.E. WILLIAMS	
	Tectonic and stratigraphic setting of the Zhengmuguan			Abstract	146
	Formation	101		Introduction	146
	The sedimentary facies and their associations in the	***		Late Proterozoic glacial and periglacial climate	146
	Zhengmuguan Formation	103		Possible explanations of low-palaeolatitude glaciation	153
	Massive diamictite facies	103		Global refrigeration	153

Frontmatter/Prelims

More information

	Contents				
	Equatorial ice-ring system	154		Whitehill Formation	198
	Geomagnetic field non-axial	154		Palaeoclimatic reconstruction	198
	Obliquity of the ecliptic > 54°	155		Depositional model for the mudrocks	200
	Other evidence for Late Proterozoic obliquity	156		Conclusions	201
	Discussion	159		Acknowledgements	202
	Conclusions	161		References	202
	Acknowledgements	161			
	References	161	15	A palaeoenvironmental study of black mudrock in the glacigenic Dwyka Group from the Boshof-Hertzogville	204
12	Isotopic signatures of carbonates associated with Sturtian (Neoproterozoic) glacial facies, central Flinders	165		region, northern part of the Karoo Basin, South Africa D.I. COLE and A.D.M. CHRISTIE	
	Ranges, South Australia			Abstract	204
	A.R. CROSSING and V.A. GOSTIN			Introduction	204
	Abstract	165		Geology	204
	Introduction	165		Depositional environment of the Dwyka Group	205
	Regional stratigraphy	165		Black mudrock	208
		167		Description	208
	Methods			Total organic carbon and organic composition	210
	Isotopic and geochemical analysis	167		Palaeoenvironment	
	Palaeoenvironmental and diagenetic interpretation	169			210
	Interpretation of ancient dolomites	170		Conclusions and regional analysis	212
	Comparison with other Neoproterozoic dolomites	170		Acknowledgements	213
	The overlying isotopic signal	173		References	213
	Geochemistry and the carbon signal	174	16	Late Paleozoic post-glacial inland sea filled by fine-	215
	Conclusions	174	10	grained turbidites: Mackellar Formation, Central	
	Acknowledgements	174		Transantarctic Mountains	
	References	174		M.F. MILLER and J.W. COLLINSON	
	Description of the state of the	176			215
13	Reactive carbonate in glacial systems: a preliminary	176		Abstract	215
	synthesis of its creation, dissolution and reincarnation			Introduction	215
	I.J. FAIRCHILD, L. BRADBY and B. SPIRO			Stratigraphic setting	215
	Abstract	176		Facies	217
	Introduction	176		Shale facies	217
	Creation of fine detrital carbonate	178		Interbedded sandstone and shale facies	217
	Chemical controls on dissolution and precipitation	179		Massive sandstone facies	220
	Weathering reactions	179		Burrowed sandstone facies	220
	Equilibrium thermodynamics	180		Diamictite facies	221
	Kinetics factors	181		Large-scale cross-stratified sandstone facies	222
	Mechanisms of precipitation	182		Large-scale channels	223
	Ripening	182		Facies associations and depositional environments	223
	Warming	183		Facies associations	223
	Freezing	183		Depositional environments	224
	Common ion effect	183		Paleocurrents	226
	Changes in CO ₂ and alkalinity	184			226
	Evaporation and transpiration	184		Paleosalinity Biogenic structures	226
	Skeletal biomineralization			Carbon:sulfur ratios	226
		184			220
	Recent examples	184		Comparison with Upper Paleozoic post-glacial	226
	Carbonate saturations of proglacial waters	184		marine sequences	226
	Processes in the ice-marginal meltout zone	186		Summary	227
	Crusts on clasts: regelation versus vadose zone			Water depth	227
	phenomena	188		Paleogeography	227
	Post-glacial transformations	189		Model for filling of the Mackellar post-glacial inland	
	Application to glacial sedimentary sections	190		sea	228
	Crusts on clasts	190		Basinal processes and environments	228
	Diamict matrix and muds	190		Shoreline processes and environments	230
	Sorted sands and gravels	190		Model	230
	Acknowledgements	190		Conclusions	230
	References	190		Appendix. Outcrop locations	231
				Acknowledgements	231
14	A Permian argillaceous syn- to post-glacial foreland	193		References	231
	sequence in the Karoo Basin, South Africa				
	J. N. J. VISSER		17	Ice scouring structures in Late Paleozoic rhythmites,	234
	Abstract	193		Paraná Basin, Brazil	
	Introduction	193		A.C. ROCHA-CAMPOS, P.R. DOS SANTOS and J.R.	
	Spatial and age relationships of the Prince Albert and			CANUTO	
	Whitehill Formations	193		Abstract	234
	Lithology	196		Introduction	234
	Prince Albert/Pietermaritzburg Formations	196		Stratigraphic setting	234



Background

Greenland

Iceberg scouring and massive diamicton formation in

the modern glacimarine environment of East

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Frontmatter/Prelims
More information

X	Contents								
		Ice scour structures	236		Icebergs and glacimarine sediments in East	• • •			
		Discussion	237		Greenland	249			
		Acknowledgements	239		Origin of massive diamicton and ice keel turbate	249			
		References	239	5 55	Criteria for distinguishing glacimarine diamictons				
	18	Soft-sediment striated surfaces and massive diamicton	241		and related facies Discussion	252 253			
		facies produced by floating ice C.M.T. WOODWORTH-LYNAS and J.A. DOWDESWELL Abstract			Summary and conclusions	256			
			241		Acknowledgements	256			
			241		References	256			
		Introduction	241	10					
		Origins of soft-sediment striated surfaces		19	Environmental evolution during the early phase of Late	260			
		Subglacial flutes	242		Proterozoic glaciation, Hunan, China				
		Ice sheet grounding line striations	242		QI RUI ZHANG				
		Surfaces formed by free-floating ice masses	242		Abstract	260			
		Scour by free-floating ice masses: ice keel scour mark			Introduction	260			
		morphology and identification criteria	243		Geological setting	260			
		Examples of ice keel scour marks and soft-sediment			Sedimentology of the Xieshuihe Formation	260			
		striated surfaces in ancient glacial sediments	243		Features of the upper contact of the Xieshuihe				

243

244

249

Formation

Acknowledgement

Discussion

References

264

266

266

266

Cambridge University Press

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Frontmatter/Prelims

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Preface

An IGCP contribution

This volume represents the final contribution of the International Geological Correlation Program Project 260 'Earth's Glacial Record'. The project was active from 1987 until 1991. It succeeded IGCP Project 24 'Quaternary Glaciations in the Northern Hemisphere', and Project 38 'Pre-Pleistocene Tillites'. Thus, its goal was to promote further research on glaciations, whatever their age and location, and to encourage geologists working in Modern or Quaternary sequences and those working on Pre-Pleistocene rocks to share their experience and different approaches. In this way we hoped to emphasize and illuminate global aspects of glacial phenomena in addition to facies description and regional problems. Participants in the project were encouraged to consider the tectonic versus climatic control on sedimentation in glacially influenced basins, the nature of the feedback between plate positioning, tectonics, and climate, and the paleoenvironmental significance and distribution of specific rock types (e.g. black shales, carbonates, iron formations, etc.) during and immediately after glacial periods.

These ambitious objectives were addressed by three subgroups that were in charge of, respectively, geodynamic setting, paleomagnetic reconstructions, and significance of specific rock types. Yearly meetings and field trips were held in various countries (10 field trips in Canada, Brazil, USA, UK, Mali), covering a vast variety of structural terranes (active and passive margins, intracratonic basins) from Proterozoic up to Modern in age. This clearly demonstrated the interest in and necessity for such a comprehensive approach among geologists. We do not pretend that all objectives were perfectly met, but aspects have been clarified and progressively more communications on global climatic or tectonic implications were presented during the successive meetings. This awareness of the importance of a global approach is certainly one of the most significant results of the project.

A second outcome of the project is a definite improvement in communication among workers on ancient and modern deposits leading to more sophisticated facies interpretation. Such interaction, as well as the idea of submitting a common IGCP project, began at the Till Mauretania 83 Symposium (Deynoux, 1985a), when Quaternary and 'paleo' geologists examined West African

Late Proterozoic glacial deposits. The final decision to submit the project, which was encouraged and supported in particular by J.C. Crowell and N.M. Chumakov, was made after an informal meeting during the 1986 International Sedimentological Congress in Canberra.

Content of the volume

The present volume complements the largely descriptive compilation made by Hambrey and Harland (1981) for IGCP 38. It is less encyclopedic and more interpretive. It reflects results of research on the IGCP 260 project themes outlined above.

One of the most puzzling climatic phenomena in the Earth's history is the possibility of glaciations at low latitudes, inferred from low paleomagnetic inclinations in many Late Proterozoic (Neoproterozoic) strata containing glacial deposits. In response to this paradox, a glacial origin for most alleged Late Proterozoic 'tillites' was denied by Schermerhorn (1974) in an excellent and somewhat provocative paper. Among his arguments was the occurrence of diamictites in active tectonic settings, implying that many diamictites may be explained better as deposited by debris-flow processes or in association with mountain glaciations, and the association of these diamictites with rocks generally indicative of warm climate, such as carbonate or iron deposits. Such an extreme position has in many cases been undermined by field evidence (Hambrey and Harland, 1981). However, Schermerhorn's paper forced reappraisal of arguments in terms of plate tectonic activity and reinforcement of those arguments which concern the glaciogenic origin of diamictite-bearing facies associations. IGCP Project 260 extended these themes by investigating the effect of tectonic setting upon the distribution and type of sedimentary facies in proven glaciogenic successions. The first nine papers of this volume address these themes.

N. Eyles and G.M. Young (p. 1) give an overview of glaciations in Earth's history and focus on the role of plate tectonic processes in the production and preservation of glaciogenic deposits. They demonstrate that strong uplift in active collisional margins or on the flanks of basins undergoing regional extension, and the resultant enhanced weathering that causes drawdown of atmospheric

xiii

xiv Preface

CO₂, provide first-order controls on glaciations. Many points that they discuss are fleshed out in papers later in the volume. The paper by P.K. Link, J.M.G. Miller, and N. Christie-Blick (p. 29) may serve as an illustration of the Eyles and Young hypothesis on the origin of glaciations. They present well-documented examples of Late Proterozoic glacial-marine sedimentation along the margin of differentially subsiding basins that developed during an episode of rifting in western North America. They provide careful descriptions of diamictite-bearing facies associations and emphasize the common occurrence of relatively deep-water facies in most of the Late Proterozoic sequences of inferred glacial origin. In contrast, J.M.G. Miller (p. 47) illustrates the merits of Schermerhorn's arguments in an excellent example of Late Proterozoic diamictitebearing facies associations in a continental rift system in eastern North America. She proposes that diamictites were deposited in a lacustrine environment under the influence of a local alpine glaciation.

The next two papers show the influence of local tectonics upon glacial sedimentation patterns during the Late Paleozoic. V. Von Brunn (p. 60) proposes a Permo-Carboniferous model in which glacial-marine deposits, including a large amount of diamictite, were formed within and on the flanks of a subsiding trough which developed over a failed rift. Depositional architecture appears largely controlled by pre-existing topography and glacially related sea-level changes. A.B. França (p. 70) presents an overview of the stratigraphy and hydrocarbon potential of the Carboniferous-Permian Itararé Group in the whole Brazilian Paraná Basin. The Itarare Group forms a continuous and thick record of a temperate glacial-marine environment in which the distribution and thickness of sedimentary units were affected by structural lineaments. This contribution is original because it concerns an economic aspect of the glacial sequences. According to structural setting and proximity of source rocks, the Itarare Group constitutes an excellent model for petroleum exploration.

Difficulty determining the origin of massive diamictites was frequently discussed at IGCP 260 meetings, J.N.J. Visser (p. 83) addresses this problem using four examples from the Permo-Carboniferous Dwyka Formation in South Africa. He proposes well-defined criteria, based on clast fabric and facies context, for the absolute identification of mechanisms of deposition of these diamictites in a glacial marine environment. Harking back to Schermerhorn (1974), Lu Songnian and Gao Zhenjia (p. 95) questioned the origin of two superposed Late Proterozoic diamictite-bearing formations in West China. Both formations were previously interpreted as glacial. Using stratigraphic and sedimentologic arguments, the authors demonstrate that the 'lower diamictite' corresponds to non-glacial debris-flows, the 'upper diamictite' to continental glacial deposits. Zheng Zhaochang and Li Yuzhen (p. 101) report Late Proterozoic glaciogenic successions in northwestern China, which consist of massive to bedded diamictite layers deposited subglacially or as subaqueous debris flows, overlain by transgressive post-glacial, thinly bedded siltstones and shales. They emphasize the importance of regional and local facies context in inferring the glaciogenic origin of the diamictite.

Lastly, the paper by M.R. Gipp (p. 109) provides sedimentological models for the large-scale architecture of glaciated shelf and slope systems in Late Cenozoic active (Gulf of Alaska) and passive (Nova Scotia) margins. Processes of deposition appear identical on both active and passive margins, but the gross depositional architecture of glacial marine deposits differs depending on the preservation potential of sediments, which is controlled by tectonics and relative sea-level changes.

Recent years have seen the introduction of the exciting new concepts of sequence stratigraphy. Such ideas are particularly relevant to glaciogenic successions because of the associated rapid and large-scale sea-level changes. However, sequence stratigraphy remains scarcely used in glacial rock successions owing probably to local effects related to the common occurrence of glacigenic deposits in tectonically active areas. J.N. Proust and M. Deynoux (p. 121) propose a sequence stratigraphic model based on the definition of a depositional genetic unit and its evolution through space and time in the marine to continental transitional zone of an intracratonic glacially influenced basin. Their model is developed from detailed field analysis of Late Proterozoic glacially related deposits on the West African platform. These genetic units and their different development in a stacking pattern lead to the definition of different orders of stacked sequences that are interpreted in terms of short-term climatically (glacially) controlled and long-term tectonically driven baselevel fluctuation cycles.

As shown by the papers quoted above, geologists have strong arguments which confirm the glacial origin of several Late Proterozoic successions. However, this does not solve the problem of low latitude glaciations (Chumakov and Elston, 1989). Hypotheses such as fast-moving plates (Crowell, 1983) or global glacial climate (Harland, 1964) have been proposed but are difficult to support on paleomagnetic and geologic grounds. Astronomical causes have also been proposed (Williams, 1975, Sheldon, 1984), and in this volume G.E. Williams (p. 146) again addresses this problem which 'challenges conventional views on the nature of the geomagnetic field, climatic zonation, and the earth's planetary dynamics in Late Proterozoic time'. Williams gives new evidence (paleomagnetic and time-series analysis of tidalites, and paleoclimatic interpretation of periglacial structures) supporting his previous hypothesis of a large obliquity of the ecliptic (>54°) leading to a reverse climatic zonation and marked seasonality.

Although still used as an argument against glaciation, the cooccurrence of carbonate rocks and glacial deposits is common. However, there are many facets of the association. Skeletal carbonates are common in high latitude seas today but they do not have exact equivalents in Proterozoic rocks. As reviewed by Fairchild (1992) the extensive ice sheets of Late Proterozoic times appear to have encroached onto previously warm carbonate-forming environments which returned following glaciation. Glacial deposits are carbonate-rich primarily because of the incorporation of detrital carbonate. Subglacial redistribution of detrital carbonate by dissolution and reprecipitation by stress-related melting-freezing processes was proposed for some Late Proterozoic terrestrial tillites (Deynoux, 1985b). The discovery of marine recrystallization of

Preface

Late Proterozoic glacially transported rock flour in Svalbard (Fairchild et al., 1989) gave new insight to the problem. Following this discovery, Crossing and Gostin (p. 165) investigated examples of diamictites in the Adelaide geosyncline of south Australia and found geochemical evidence from the composition of the matrix for deposition in a sea diluted by meltwater. Additionally, correlation between high Fe and reduced δ^{13} C values in the matrix suggests that bacterial activity accompanied the diagenesis of the rock flour. In order to gain a better understanding of the chemical processes involved when carbonate rock flours interact with fluids in glacial systems, investigations of carbonate-rich Quaternary glacial systems have been started, and I.J. Fairchild, L. Bradby, and B. Spiro (p. 176) report their preliminary conclusions. Although they stress the importance of postglacial processes in controlling lithification of the sediments, evidence is also found for the precipitation of calcite in the matrix of a refrozen meltout till. Since this article was prepared they have found similar material in unfrozen meltout till and basal ice.

The succession of glacial deposits in Late Paleozoic sequences across much of Gondwana is punctuated by a sharp contract between diamictites and overlying late- or post-glacial dark- to black-colored mudstones (Domack et al., 1992). Such a sharp contact has also been described in intracratonic Late Proterozoic and Late Ordovician glacial sequences on the West African platform (see Deynoux and Trompette in Hambrey and Harland, 1981). The diamictite-black mudstone transition is important in that it is widely believed to represent the end of glacial climates within the regional extent of the basins in which it is found. Hence these rocks preserve a record of global warming associated with an apparently rapid 'glacial' to 'interglacial' transition. A similar hypothesis was also proposed for certain Proterozoic carbonate horizons capping glacial deposits (e.g. Williams, 1979). Depositional mechanisms for the dark to black mudstones are varied. J.N.J. Visser (p. 193) describes the conditions that prevailed in a shallow to moderately deep Late Carboniferous foreland basin (Dwyka glaciations of the Karoo Basin) after self-destructive collapse of a marine ice sheet resulting from a relative sea-level rise. The syn- to post-glacial dark to black mudrocks, which overlie the glaciogenic deposits, were deposited by suspension settling of mud and mud turbidites. Basin tectonics, oceanic circulation, and climate controlled the organic-rich black mud deposition. D.I. Cole and A.D.M. Christie (p. 204) also describe Early Permian black mudrocks overlying diamictites (debris and turbidity flows) deposited by a retreating tidewater glacier during the final phase of the Dwyka glaciation. The black mudrocks are the product of pelagic mud settling proximal to the ice front where freshwater plumes mixed with basinal saline water. They account for a sudden increase in organic productivity, indicating that the rapid termination of the Dwyka glaciation was accompanied by a sharp rise in temperature. M.F. Miller and J.W. Collinson (p. 215) describe the processes and environments that characterize the filling of a large Lower Permian post-glacial inland sea in Antarctica. Deposition in relatively shallow water was dominated by turbidity currents carrying finegrained sediments in channel-overbank systems. The glacial environment allows the definition of a model in which a finegrained turbidite system is paradoxically fed by a coarse-grained braided stream of the outwash plain.

Recent years have seen increasing interest in the sedimentary record of ice scours on continental shelves and lakes. Such structures and associated diamict deposits are widespread on Pleistocene shelves but are rarely known from the rock record. The paper by A.C. Rocha-Campos, P.R. dos Santos, and J.R. Canuto (p. 234) describes Early Permian iceberg scours. These scours are encountered on bedding planes of rhythmites inferred to be varves in a relatively deep freshwater body. C.M.T. Woodworth-Lynas and J.A. Dowdeswell (p. 241) argue that many striated surfaces, and associated diamictites found in ancient successions may have been produced by marine (or lacustrine) ice keel scour and report a modern analogue from the Greenland shelf. The need for a more critical examination of ancient glacial striated surfaces and associated facies is clearly indicated. Qui Rui Zhang (p. 260) describes periglacial indicators such as iceberg scours and dropstones, ice wedge casts, and glaciotectonic structures. These structures indicate that, instead of an abrupt erosional unconformity, the glaciogenic deposits of the Nanhua Ice Age (Late Proterozoic) of South China are locally conformably underlain by rocks that mark a progressive climatic transition from warm to cold.

Significance

The results of IGCP Project 260 are very relevant to current concerns about global change. As more details of the present climate and the climate of the Pleistocene are revealed, the interplay of ocean, atmosphere, biosphere and lithosphere with the internal heat engine of the Earth is being revealed as something of great complexity. We are still far from a complete understanding of the Earth's climatic system but glaciogenic deposits provide critical palaeoclimatic data. In spite of recent suggestions that some tillites/ diamictites may be the ejecta of large planetesimal impacts (Rampino, 1992; Oberbeck et al., 1993), extensive research and field studies demonstrate that the vast majority of documented ancient glaciogenic deposits correctly record periods of cold climate during Earth history. The discovery of mechanisms governing the appearance and disappearance of glaciers on Earth is paramount to the understanding of long-term climatic change, and the only longterm record of climatic change is the geologic record. The rock record of glaciation has been catalogued (Hambrey and Harland, 1981) but still has not been perfectly described. Since the 1983 Mauritanian meeting (Deynoux, 1985a), good facies descriptions of ancient glacial deposits have been proposed, and lately sequence stratigraphic concepts have developed. Now glacial sedimentology has joined the mainstream of sedimentology. In the near future, because most ancient glacial evidence is preserved in marine sequences, we must achieve a better understanding of the workings of glacially related marine basins.

xvi Preface

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xvii